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Synergy of multispectral and multisensors satellite observations to evaluate desert aerosol transport and impact of dust deposition on inland waters: study case of Lake Garda

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Abstract. The capabilities of different Earth Observation multispectral satellites are employed for detecting and tracking of desert dust coming from North Africa toward the Northern Italy area and for evaluating the impact of Saharan dust deposition in inland waters, such as those of Lake Garda. Absorbing and scattering spectral optical properties of desert aerosol in the atmospheric windows in the ultraviolet, visible-near-infrared, and infrared spectral ranges are exploited in the dust retrieval performed by OMI/Aura, MODIS/Terra-Aqua, and SEVIRI/MSG satellite sensors. Therefore, the direct link between dust deposition and increase in phytoplankton abundance has been assessed retrieving MERIS-based chlorophyll-a (chl-a) concentration for the desert dust events. Estimates of the increased chl-a in the lake have been derived with values in concentration from 30% to 170%. AERONET sun-photometer measurements, gravimetric particulate matter samplings, *in situ* chl-a concentration and surface temperature are employed to select events and assess the presence of desert dust and recognize a corresponding increase of the phytoplankton abundance in the analyzed inland waters. The improved observational features that will be provided by the next European Sentinels missions, namely Sentinel-2, 3, 4, 5P, together with MTG-I (Imager) and MTG-S (Sounder) will allow better monitoring atmospheric constituents and studying the environmental impacts of desert dust transport. © The Authors. Published by SPIE under a Creative Commons Attribution 3.0 Unported License. Distribution or reproduction of this work in whole or in part requires full attribution of the original publication, including its DOI. [DOI: [10.1117/1.JRS.9.095980](https://doi.org/10.1117/1.JRS.9.095980)]

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1 Introduction

Aerosol particles play an important role in atmospheric processes affecting climate, air quality, rain formation, biogeochemical cycles, representing a challenging subject of analysis in climate research, weather forecast, and, in particular, in the remote sensing retrieval of surface properties. Among the various aerosol types, mineral dust can be considered one of the major sources of tropospheric aerosol loading, and it has been extensively studied for its impact on the radiative balance by the absorption and scattering of sunlight^{1,2} and on the hydrological cycle by means of their role as cloud condensation nuclei.³ The optical and physical properties of dust aerosol depend on the different sources and the changes they undergo to during their transport in the atmosphere. Dust aerosols are blown into the atmosphere mainly from arid regions and are subject to long-range transport at an intercontinental scale. Among the main dust transport patterns, it is worth mentioning those from North Africa over the Atlantic Ocean, from Arabian Peninsula toward Arabian Sea and Indian Ocean, and from East Asia over the Pacific Ocean.

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Concerning the European areas, the major transport patterns of Saharan dust are those originating in the North African toward the Southern European regions.⁴

Desert dust transport events can also affect air quality and its management. The European Commission (EC) directive of 2008 on “Ambient air quality and cleaner air for Europe”⁵ allows to subtract the number of occurrences of PM₁₀ concentration exceedances above the daily limit of 50 $\mu\text{g}/\text{m}^3$ when due to natural source. Thus, an increased attention to monitor the dust transport, also by means of remote sensing,^{6,7} is hence required.

Moreover, the Saharan dust deposition in the oligotrophic Mediterranean Sea is a source of nutrients for phytoplankton due to the presence of trace metal as Fe, Al, and Si.^{8,9} As a typical feature, due to the presence of iron oxides (hematite) and phosphorus, which mainly in oligotrophic stratified waters could be limiting factors for phytoplankton growth, dust depositing on water surface could affect the abundance of phytoplankton,¹⁰ making nutrients available for primary production. Several studies in very low trophic waters also showed summer algal blooms as response to N and P depositions after dust events.^{11,12} This effect has been successfully investigated using satellite observations in the Mediterranean Sea by Refs. 13–17, which allow a frequent retrieval of water quality parameters and thus analyzing time pattern with a synoptic vision on the entire case study area.

Aerosol particles, being both of natural or anthropogenic origin, are characterized by optical properties varying in the spectral range from visible (VIS) to near-infrared (NIR) wavelengths.^{18–20} Desert dust aerosol displays the strongest absorption features in the ultraviolet (UV) through the VIS wavelengths,¹⁸ with respect to other aerosol types. This behavior is due to dust mineralogical composition (mainly hematite and silica) affecting dust radiative properties also in the thermal infrared (IR) spectral region. In particular, the presence of clays and quartz in the mineral composition increases the sensitivity of the IR spectral brightness temperature (BT) to desert dust.^{19,20} Thus, in order to highlight the presence of desert dust in the atmosphere, the spectral features from the UV to the thermal-IR can be profitably exploited using observations carried out by optical satellite sensors.

Among active spaceborne instruments, lidar devoted to aerosols and clouds observation such as Cloud-Aerosol Lidar with Orthogonal Polarization (CALIOP), on board the Cloud-Aerosol Lidar and Infrared Pathfinder Satellite Observations (CALIPSO) satellite, which is a joint program of National Aeronautics and Space Administration/NASA and Centre National d’Etudes Spatiales/CNES, furnishes vertical profile of aerosols and clouds optical properties. Respect to the passive instruments, the main advantages presented by CALIOP are the capability to discriminate between dust and other types of aerosol by means of the depolarization measurements at 532 nm and the possibility to perform measurements both during day and night. However, due to its repetition cycle of 16 days, CALIOP observations are limited in temporal coverage and, then, mainly adopted for a seasonal climatology of the desert aerosol.²¹

The objectives of this study are twofold. At first, a synergistic use of multispectral remote sensing observation acquired by different satellite passive sensors is applied to detect and track the transport of desert dust toward the Lake Garda area, i.e., the study area located in Northern Italy, and then the effect of the identified dust deposition in the lake is estimated through the monitoring of the trend of chlorophyll-a (chl-a) concentration as proxy of phytoplankton abundance.

The first aim is reached by (i) analyzing the brightness temperature differences (BTD) between 8 and 11 μm channels versus BTD between 11 and 12 μm channels coming from radiance measured by the EUMETSAT/MSG Spinning Enhanced Visible and Infrared Imager (SEVIRI) with a 15 min repetition cycle; (ii) employing the spectral trend of aerosol optical properties (AOP) retrieved from the observations of NASA/Aura Ozone Monitoring Instrument (OMI) in the spectral range (300 to 500) nm; and (iii) exploiting the high capability of NASA/Terra, Aqua Moderate Resolution Imaging Spectroradiometer (MODIS) to retrieve the dust optical properties in the VIS–NIR spectral range and discriminating the presence of clouds mainly thanks to spectra simultaneously registered in the IR channels.

The second aim is reached by processing the ESA/Envisat Medium Resolution Imaging Spectrometer (MERIS) full resolution (FR) level 1 data to retrieve chl-a concentration values, at spatial resolution of 300 m, in the time window of each event analyzed and by completing

the analysis with ancillary *in situ* measurements, which allows validating the actual effect of dust deposition.

The synergistic approach adopted in this study is finally analyzed in view of the future European sentinels missions, namely Sentinel 3 and 5P, combined with MTG-I (Imager) and MTG-S (the UV-VIS-NIR Sounder, named UVN, boarding the Sentinel 4), which would be successfully and properly utilized in the next years.

2 Materials and Methods

2.1 Study Area and Case Studies Selection

This study has been focused on Lake Garda but for taking into account desert dust transport events, which are characterized by typical large geographical pattern, *in situ* data have been collected on a wider area (Fig. 1): from Bormio (46.5°N; 10.4°E) to Modena (44.6°N; 10.9° E), in the north–south direction, and from Ispra (45.8°N; 8.6°E) to Venice (45.4°N; 12.3°E) in west–east direction.

Lake Garda is the largest lake in Italy (368 km²), on average 133-m deep, located at 65 m above sea level in the subalpine area in Northern Italy. It is a glacial oligo-mesotrophic oligomictic lake; phosphorus concentration in the epilimnion is comprised between 4 and 10 mg/m³, chl-a mean concentration²² varies between 0.5 and 12 mg/m³, and it has a high residence time (26.6 years).²³ A strong variability affects chl-a trend during the year; higher values are recorded between the end of the winter and the beginning of spring (and in some cases in autumn), when algal blooms mainly occur, while more stable conditions are recorded in summer season, with lower concentration values.^{24–26}

Complete circulation, which brings up nutrients from the bottom, does not occur every year; during MERIS operational lifetime it was recorded in 2004, 2005, and 2006.²⁷

The four case studies analyzed in this work, parted in dust and no-dust events and also representative of different seasons, are presented in Table 1 together with the corresponding satellite- and ground-based data available.

Time series of PM₁₀ and PM_{2.5} (particulate matter with aerodynamic diameter less than 2.5 and 10 μm, respectively) concentrations were used to single out dust events. In particular, the daily data were sampled in Bormio (alpine site), Bosco Fontana (B. Fontana, rural site), and Ponti sul Mincio (Ponti S.M., urban site). Together with these data, aerosol optical depth (AOD)



Fig. 1 The study area. The dark area inside the white box is the Lake Garda. Chl-a was analyzed in the area highlighted in the small red box over the lake (MERIS image). Red symbols indicate AERONET sites; green symbols indicate ARPA Lombardy meteorological stations.

Table 1 Case studies in the domain of interest analyzed in this study, parted in “dust” and “no-dust” events, together with satellite- and ground-based observations available in the different seasons/periods. Ispra, Modena, and Venice refer to AERONET sites. For the MERIS/Envisat observations, useful not cloudy days relative to satellite overpass are indicated. In the event (iii) useful MERIS data were not available (see text).

| Event—season—period | EO satellite sensor—observation feature | | | | Ground-based observation—site | | |
|---|--|-----------------------------------|-----------------------------------|---|---|-----------------------|--------------------|
| | SEVIRI | OMI | MODIS | MERIS | PM samplings | AOD, α | Precipitation |
| Dust events | | | | | | | |
| (i)—Summer—July 27 to 29, 2005 | Data over the whole period, every 30 min | Daily acquisition over the period | Daily acquisition over the period | June 24 July 12, 20, 23, 26, 29 August 1, 4, 8, 17, 24, 30 September 2 | PM ₁₀ : Bormio, Ponti S.M., B. Fontana PM _{2.5} : Ponti S.M., B. Fontana | Ispra, Modena, Venice | Bormio, Ponti S.M. |
| (ii)—autumn—October 4 to 10, 2007 | Data over the whole period, every 30 min | Daily acquisition over the period | Daily acquisition over the period | September 22 October 2, 8, 15, 21 November 3 | PM ₁₀ : Bormio, Ponti S.M., B. Fontana PM _{2.5} : Ponti S.M., B. Fontana | Ispra, Modena, Venice | Bormio, Ponti S.M. |
| No-dust events | | | | | | | |
| (iii)—summer—June 1 to 14, 2005 | Data over the whole period, every 30 min | Daily acquisition over the period | Daily acquisition over the period | — | PM ₁₀ : Bormio, Ponti S.M., B. Fontana PM _{2.5} : Ponti S.M., B. Fontana | Ispra, Modena, Venice | Bormio, Ponti S.M. |
| (iv)—winter—October 30 to November 30, 2006 | Data over the whole period, every 30 min | Daily acquisition over the period | Daily acquisition over the period | October 30, November 2, 3, 8, 12, 21, 30 | PM ₁₀ : Bormio, B. Fontana PM _{2.5} : B. Fontana | Ispra, Venice | Bormio, Ponti S.M. |

at 440 nm and Ångström parameter (α), calculated between 440 and 870 nm, measured at the AERONET stations in Ispra, Modena, and Venice (Fig. 1) was analyzed to recognize the presence of desert aerosol being α representative of the suspended aerosol particles size distribution. Values around or greater than 2 indicate the presence of fine particles, mainly originated by anthropogenic activities and biomass burning, while values lower than or near 1 indicate the presence of coarse particles such as desert dust; intermediate values indicate a mix of both sizes.¹⁸

Evidences of desert dust event occurrence were given by the concomitance of high values of daily samplings of PM_{10} (e.g., daily values exceeding the monthly average value added to its corresponding standard deviation), of high values (with the same criteria used for PM_{10}) of AOD (typically greater than 0.4) as well as by low values of α (typically near 1). Table 1 shows the two dust events singled out based on these criteria: event (i) July 27 to July 29, 2005 and event (ii) October 4 to October 10, 2007 as reported in Table 1. On July 28, 2005, day (day of the year) 209, PM_{10} and $PM_{2.5}$ in B. Fontana increased reaching the maximum values of 66 and 28 $\mu\text{g}/\text{m}^3$, respectively. Simultaneously, AOD measured in Ispra (Modena, Venice) site was 1.77 (0.53, 0.50), being the relative threshold equal to 0.61 (0.35, 0.38), with a corresponding α value of 0.8 (0.95, 1.0).

From October 4 to October 10, 2007 (day from 277 to 283), in Ponti S.M., PM_{10} and $PM_{2.5}$ increased to 80 and 40 $\mu\text{g}/\text{m}^3$ (20% over the relative threshold), respectively, while in Modena (Ispra, Venice) site AOD and α were around 1.03 (0.41, 0.59), being the relative thresholds equal to 0.61 (0.35, 0.31) and 0.49 (1.03, 1.10), respectively.

In the analysis of the aerosol behavior, also a no-dust event [event (iii) in Table 1] related to June 1 to June 14, 2005 period was analyzed. Beside the absence of transported dust, this event is also characterized by a relatively small presence of anthropogenic pollution. In Ponti S.M., the average over the period for PM_{10} was 50 $\mu\text{g}/\text{m}^3$ with a standard deviation equal to about 5 $\mu\text{g}/\text{m}^3$, resulting below the annual mean value of 65 $\mu\text{g}/\text{m}^3$. The mean values of AOD and α over the period at Ispra (Modena, Venice) were 0.42 (0.32, 0.28) and 1.59 (1.62, 1.78), respectively. Concerning this no-dust event the analysis of satellite observations has been performed in terms of UV–VIS–NIR AOP, as cloud covers hindered the use of MERIS to investigate the chl-a dynamics.

To minimize the risk of erroneous interpretation of the effect of dust on phytoplankton growth, a second no-dust event relative to a period during which AOD and PM concentrations increased due to anthropogenic activities was selected: i.e., event (iv) of Table 1, October 30 to November 30, 2006. In this period, AERONET data were not available for Modena but in Ispra (Venice) site AOD and α presented monthly averaged values equal to 0.15 (0.45) and 1.33 (1.56), respectively. Data from Ponti S.M. were not available but in Bormio the average PM_{10} was 19 $\mu\text{g}/\text{m}^3$ with standard deviation of 6 $\mu\text{g}/\text{m}^3$. On November 11, 14, 15, 16, 17, 23, 24, and 25, 2006 (day: 315, 318, 319, 320, 321, 327, 328, 329) PM_{10} concentrations were above the threshold (25 $\mu\text{g}/\text{m}^3$), being 46, 32, 34, 27, 28, 26, 26, and 29 $\mu\text{g}/\text{m}^3$, respectively. For this reason, analysis on phytoplankton abundance was carried out around this time window.

2.2 Satellite Data

Desert dust transport events are very often characterized by very large plumes traceable with Earth Observation (EO) satellites. In this context, EO imagery was used to track Saharan dust plumes traveling toward Italian coasts and up to register the presence of dust over Northern Italy and, in particular, the Lake Garda. For this purpose, in the next sections, we consider data analysis techniques using IR observations coming from EUMETSAT MSG geostationary platform, and UV–VIS–NIR observations coming from daily observations from satellite sensors on board NASA Aura, Aqua, and Terra Sun-synchronous platforms. AOP level 2 data from OMI and MODIS observations have been adopted to highlight the presence of desert dust in the atmospheric column and to derive a level 3 data of satellite-based map of PM concentrations at surface, respectively.

2.2.1 SEVIRI/MSG IR spectra analysis

The detection of desert dust is based on the analysis of the radiance sampled by SEVIRI on board MSG-1 (Meteosat-8) meteorological satellite²⁸ with capability to observe the high temporal

frequency and also in the night. IR observations in the atmospheric windows between 8 and 12 μm have been exploited considering the different spectral trend of desert dust optical properties with respect to those of clouds and in clear conditions (i.e., without the presence of dust or clouds). In particular, BTs are calculated from the SEVIRI level 1.5 data in the 8.7, 10.8, and 12.0 μm channels.²⁹ Then, two BTDs are derived as $(\text{BT}_{12.0} \text{ to } \text{BT}_{10.8})$ and $(\text{BT}_{10.8} \text{ to } \text{BT}_{8.7})$ enabling the enhancement of the presence of desert dust [Fig. 2(a)].

2.2.2 UV–VIS aerosol optical properties and OMI/Aura observations

Desert AOP as single-scattering albedo (SSA) and AOD present spectral trends that can be profitably used to highlight the presence of desert dust.^{18,19} Within the UV–VIS spectral range, SSA relative to this type of particles increases with wavelength. Accordingly, the absorbing AOD (AAOD) decreases while the extinction AOD (EAOD) typically decreases slowly with respect to the other types of atmospheric aerosols. For seven different types of aerosols, Fig. 3 shows the spectral trend of the SSA and the normalized EAOD, EAOD_n

$$\text{EAOD}_n(\lambda) = \frac{\text{EAOD}(\lambda)}{\text{EAOD}(550 \text{ nm})} = \frac{k_e(\lambda)}{k_e(550 \text{ nm})} \quad (1)$$

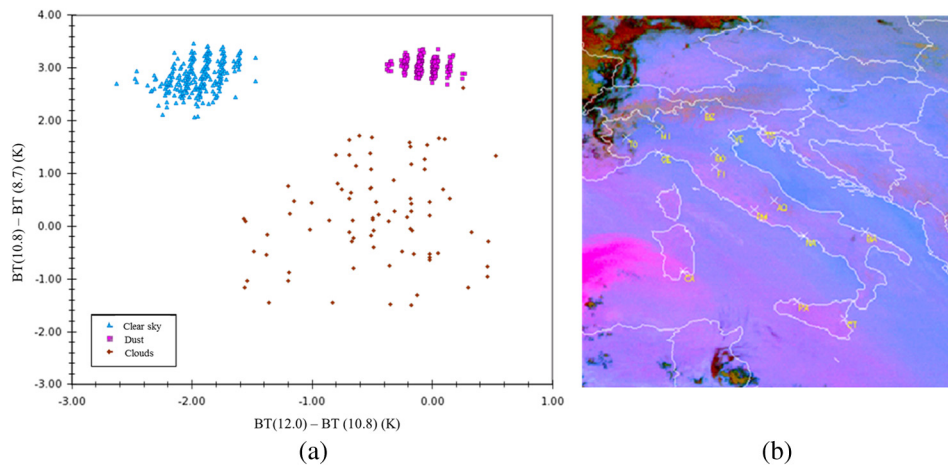


Fig. 2 (a) Brightness temperature differences plot discriminating the presence of desert dust (solid squares in magenta) and (b) composite false color RGB image (see also Sec. 3.1.1), relative to IR SEVIRI observation at 8.7, 10.8, and 12 μm on July 27, 2005, 01:00 UTC. Data refer to the event (i) in Table 1.

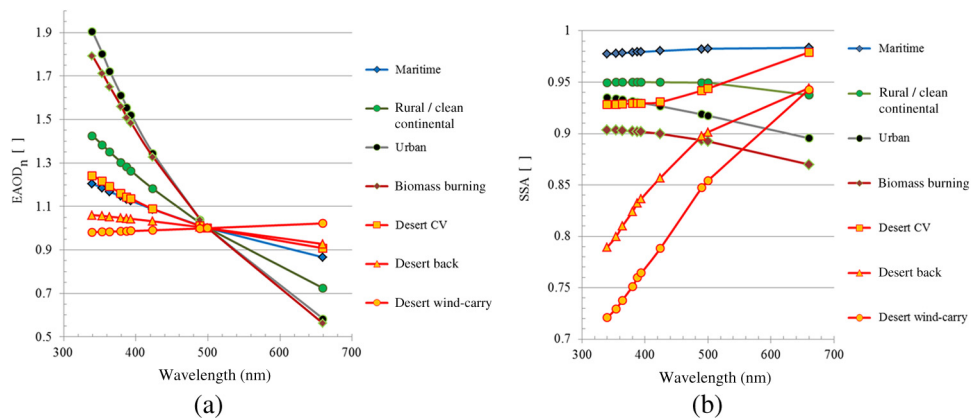


Fig. 3 (a) Spectral trend of normalized extinction aerosol optical depth, Eq. (1) and (b) single-scattering albedo derived from Ref. 30 for three desert dust and for other aerosol types.

being k_e ($\text{km}^{-1} \text{cm}^{-3}$) the aerosol-specific extinction coefficient derived from the AOP DB of Ref. 30. The aerosol microphysical properties needed to compute the optical properties of selected aerosol types of Fig. 3 are taken from Ref. 19 for the desert CV, urban, and biomass burning aerosols, from Ref. 31 for the desert background (desert back) and desert wind-carry aerosols, and from Ref. 32 for the rural/clean continental and maritime aerosols. As it can be noticed, in Fig. 3(a), excepted for the maritime aerosol, the three desert aerosol types can be discriminated with respect to the other aerosol classes.

By considering the optical absorbing features [i.e., the SSA plot of Fig. 3(b)], the significant increasing spectral trend of the desert aerosols SSA with respect to the other aerosol types can also confirm the dust presence.

UV–VIS spectra registered by OMI sensor onboard the Finnish–Dutch EOS Aura spacecraft are usefully employed. A spectral analysis of the AOP (level 2G gridded data products, OMAERUVG.003) from OMI/Aura data products (15 orbits/day; $13 \times 24 \text{ km}^2$ spatial resolution at nadir) into $0.25 \text{ deg} \times 0.25 \text{ deg}$ global grids products have been carried out to recognize the presence of the dust over the study area for the selected periods (cf. Table 1).

2.2.3 MODIS satellite-based PM using VIS–NIR observations

A Northern Italy synoptic view of $\text{PM}_{2.5}$ and PM_{10} concentrations is reconstructed³³ from MODIS/Terra and Aqua aerosol level 2 data (C005 collections) at $10 \times 10 \text{ km}^2$ spatial resolution, meteorological mixing layer, relative humidity fields simulated by MM5³⁴ model at spatial resolution of $12 \times 12 \text{ km}^2$, and *in situ* $\text{PM}_{2.5}$ and PM_{10} as described in Ref. 35.

MODIS-based PM data (Fig. 4) are here selected to compute maps of $\text{PM}_{2.5}$ to PM_{10} concentrations ratio averaged over the period of the selected events (Table 1). This ratio assumes typical values around or higher than 0.75 in case of anthropogenic pollution and around or lower than 0.6 in highlighting the presence of coarse particles, such as desert dust aerosol.³⁷

2.2.4 Retrieval of chl-a concentrations from MERIS data

To assess the effect of dust deposition on Lake Garda waters, 26 MERIS level 1B FR (300 m) images relative to years 2005 to 2007 (Table 1) were used. The images were processed through the ESA-BEAM toolbox, in order to remove radiometric disturbances as well as atmospheric and adjacency effects. In particular, the BEAM Smile correction was applied to all images.³⁸ The Smile corrected images were processed with the BEAM plug-in Improved Contrast between Ocean and Land (ICOL), to remove the adjacency effects of land on water.³⁹ To retrieve chl-a concentrations, the MERIS case 2 water BEAM processor (C2R)⁴⁰ was used as it was successfully adopted to produce MERIS-derived chl-a concentration in Lake Garda.^{41,42} MERIS level-1 and C2R level-2 flags were used for excluding pixels with risk of clouds and glint. The trend of chl-a concentration was finally evaluated in a pelagic area (about 17.5 km^2 in Fig. 1) as it is most likely the least affected by residual adjacency effects.

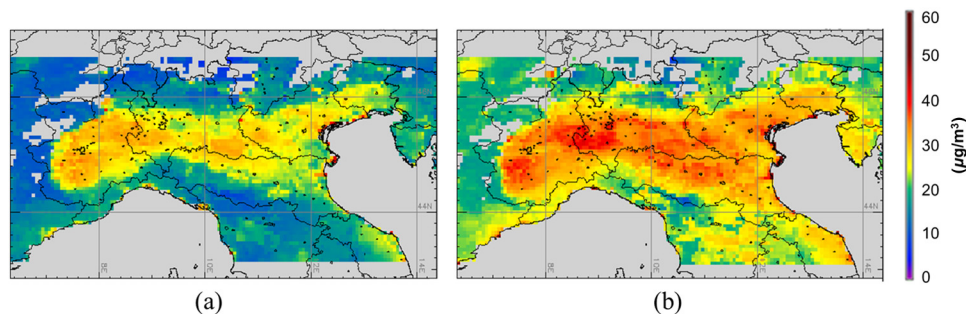


Fig. 4 2008 yearly average of MODIS-based particulate matter concentrations at surface over Northern Italy: (a) $\text{PM}_{2.5}$ map and (b) PM_{10} map. Reproduced with permission, courtesy of Alessandra Cacciari.³⁶

3 Results and Discussion

3.1 Detecting Desert Dust and Retrieval of Desert Optical Properties Using Satellite-Based Observations

In this section, the results for the four case studies (Sec. 2.1) according to remote sensing techniques presented in Sec. 2.2 are reported.

3.1.1 SEVIRI/MSG dust tracking

As introduced in Sec. 2.2.1, BTDs analysis of the SEVIRI data has been applied to the selected events (Table 1) for tracking and evaluating the origin of the desert dust plume.

Sequences of RGB (red, green, blue) composite images ($R = BT_{12.0} - BT_{10.8}$; $G = BT_{10.8} - BT_{8.7}$; $B = BT_{10.8}$) are produced using the SEVIRI/MSG BTD at spatial and temporal resolutions of 4 km and 15 min, respectively.

SEVIRI BTDs showed a significant transport of dust over the Mediterranean area and toward Italy for the dust events (i) and (ii) (cf. Table 1). In Fig. 2(b) desert dust plume over Mediterranean Sea toward the Italian coasts appears pink or magenta, while thick high-level clouds are red-brown, thin high-level clouds appear very dark, and clear sky pixels are blue.^{43,44} As the plume approaches Italy and also considering a decrease of particle concentrations during its transport, the dust comes less distinguishable.

3.1.2 Aerosol optical properties derived from OMI and MODIS-based $PM_{2.5}$ to PM_{10} ratio

Figure 5 presents 3-days averaged maps of three spectral values of AOP relative to the first dust event (cf. Table 1 and Fig. 2). The averaged values of EAOD and SSA are shown at 354, 388, and 500 nm. As typical properties of desert dust aerosol, the increasing spectral trend of SSA values can be clearly recognized in the values presented by the pixels located between the two main islands of Italy (southwestern part of the country). Within the desert plume over the sea, SSA values averaged in the Mediterranean area (from 39.3°N–8.3°E to 37.3°N–13.6°E) turn to be

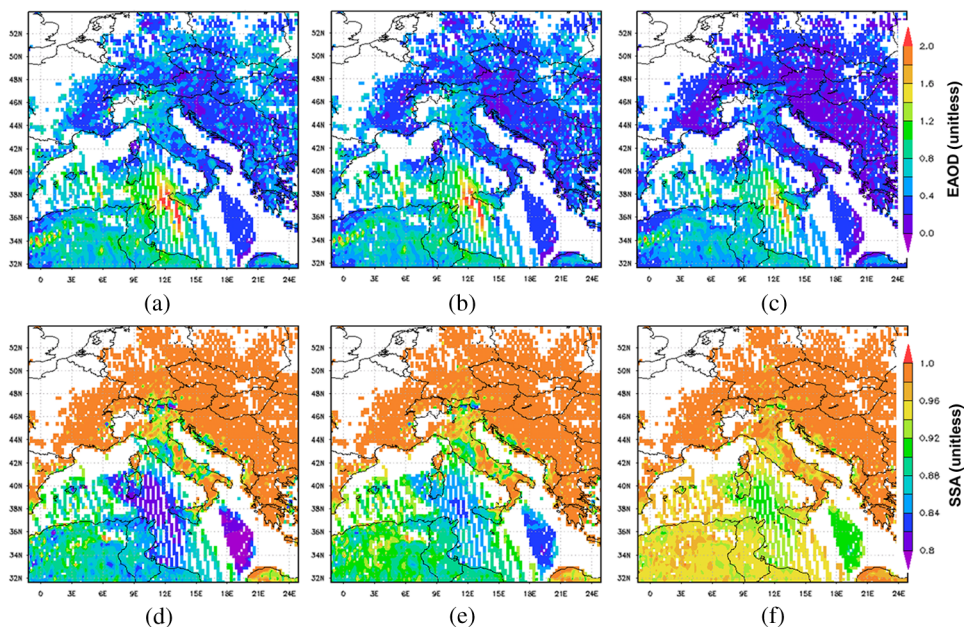


Fig. 5 UV–VIS desert dust detection over Mediterranean Sea as retrieved from OMI/Aura satellite data processing. All maps show 3-days averaged values from July 27 to July 29, 2005 of (a–c) the extinction AOD (EAOD) and (d–f) single-scattering albedo (SSA) at the wavelengths of (a, d) 354, (b, e) 388, and (c, f) 500 nm.

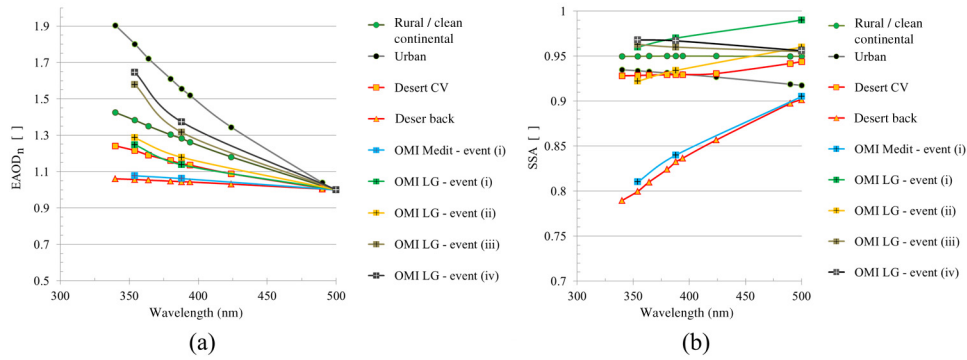


Fig. 6 Comparison between aerosol models of Fig. 3 (rural/clean continental, urban, desert CV, and desert back) and OMI-retrieved spectral averaged AOP (squared symbols) over Mediterranean (Fig. 5) and over Lake Garda areas (Fig. 7), for the event (i), and for Lake Garda area for events (ii, iii, and iv); (a) spectral trend of EAOD_n and (b) SSA values. Mediterranean and LG-event indicate Mediterranean and Lake Garda areas, respectively.

0.81 ± 0.02, 0.84 ± 0.02, and 0.91 ± 0.02 at 354, 388, and 500 nm wavelengths, respectively.

These SSA values and the corresponding spectral EAOD_n values, calculated averaging over the same Mediterranean area the spectral EAOD (Fig. 5, first row) and using Eq. (1), are plotted in Fig. 6 for a comparison with the relative spectral AOP showed in Fig. 3. It can be noticed that there is very good agreement of the retrieved OMI values (blue square) with respect to the desert background aerosol model (red triangle).

Since the study area is Lake Garda, a zoom in Northern Italy of maps in Fig. 5 is showed in Fig. 7. In the last row of Fig. 7, the OMI SSA of Lake Garda presents averaged values of 0.96 ± 0.02, 0.97 ± 0.02, and 0.99 ± 0.02 at 354, 388, and 500 nm, respectively. The mapped AOP spectral values confirm that the desert dust outbreak at the end of July 2005 reached also Northern Italy. These SSA values and the corresponding spectral EAOD_n [Eq. (1)], calculated averaging over the Lake Garda area the spectral EAOD of Fig. 7, are also included in Fig. 6

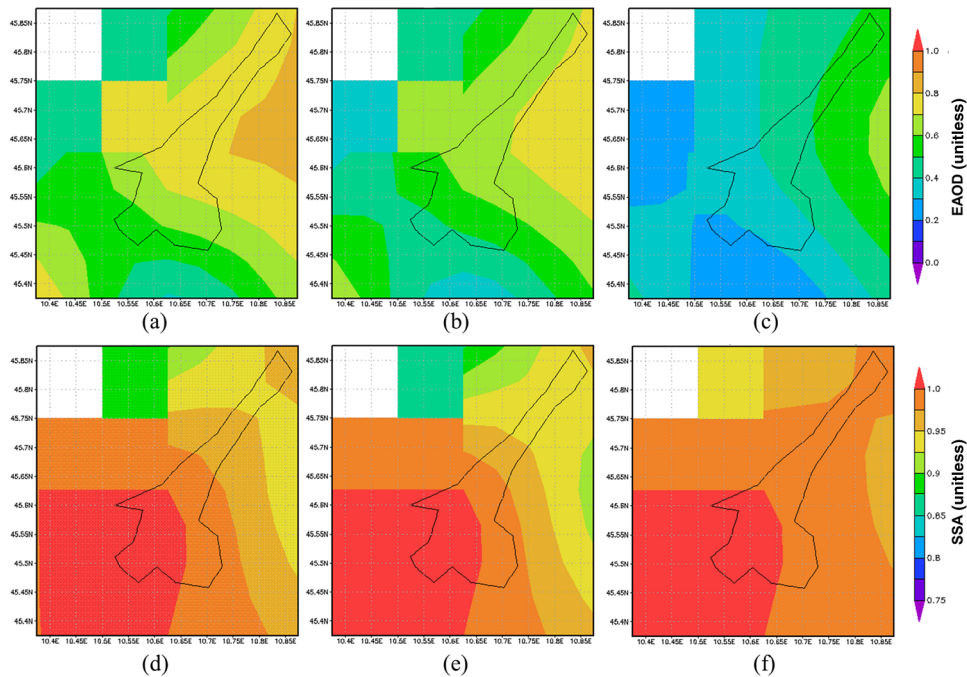


Fig. 7 Zoom over Northern Italy (Lake Garda area) of the same period of averaged maps presented in Fig 5. (a–c) the extinction AOD (EAOD) and (d–f) SSA at the wavelengths of (a, d) 354, (b, e) 388, and (c, f) 500 nm. Data refer to the dust event (i) of Table 1.

(squared green symbols) to compare them with the modeled spectral AOP depending on aerosol types rural/clean continental, urban, desert CV, and desert background.

In this event, the spectral trend of retrieved OMI AOP (EAOD_n and SSA) confirms the presence of a significant component of dust in the study area. Figure 6 shows that the increasing spectral trend of OMI SSA and EAOD_n is very similar to those ones of desert CV aerosol model (square red symbols, yellow filled). Nonetheless, the higher values of retrieved OMI SSA and the slight difference of retrieved EAOD_n with respect to the SSA and EAOD_n values, respectively, of desert CV aerosol type can be justified considering a proper mixing with a more scattering aerosol type as the rural. In particular, for these hydrophilic aerosol particles, high values of the relative humidity also contribute to increase the scattering features of the aerosol particles.⁴⁵ To provide an exhaustive picture of the desert dust case study, a selection of satellite-based information is presented in Fig. 8. In particular, NO₂ tropospheric column amount is shown [Fig. 8(a)] as derived from daily OMI data processing⁴⁶ and averaged over the period of the event (i), July 23 to July 31, 2005. The values highlight a relatively small contribution of human activities to air pollution around the study area.

MODIS-based PM_{2.5} and PM₁₀ daily concentrations for the investigated periods (cf. Table 1) have been used to compute maps of the MODIS-based PM_{2.5} to PM₁₀ ratio and then averaged over the dust and no-dust periods.

An example of this map averaged over the July 23 to July 31, 2005 time period comprising the first event (i) and referred to the Lake Garda area is reported in Fig. 8(b). The average value of the PM_{2.5}/PM₁₀ ratio over Lake Garda is 0.61 ± 0.05 , highlighting a significant presence of coarse particles and thus confirming the presence of dust transport in the study area. This analysis applied to the entire satellite data for detecting the desert dust transport has been applied also to the second case study—event (ii)—listed in Table 1.

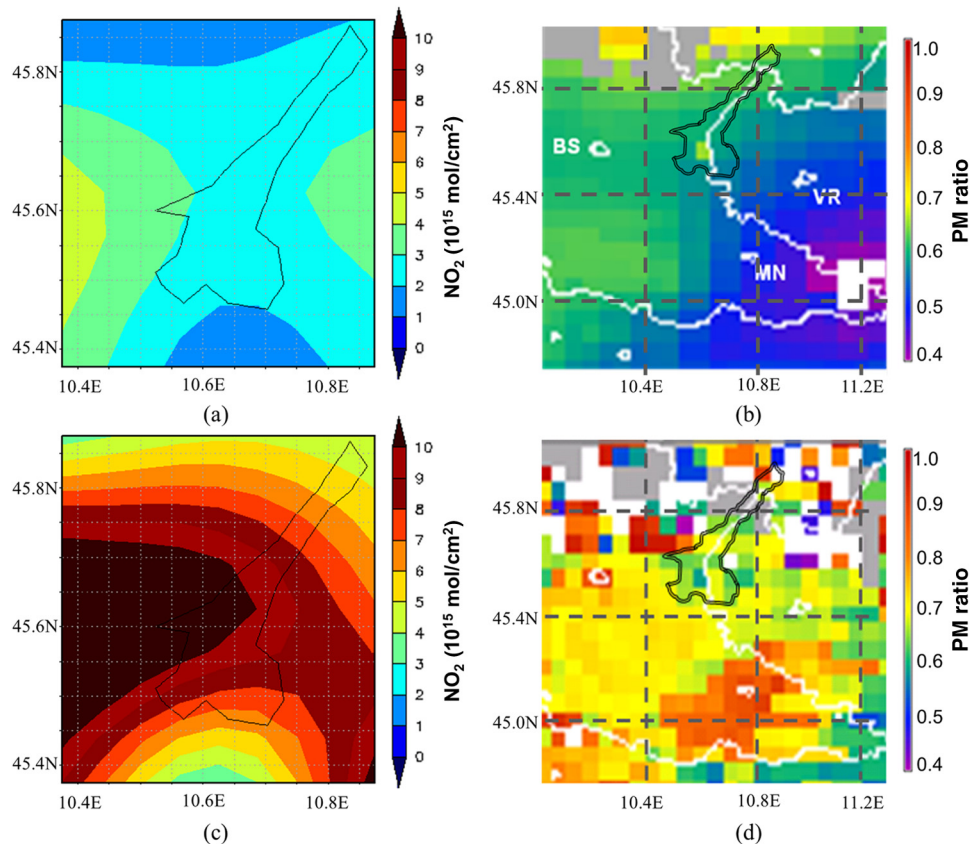


Fig. 8 On the top: average of period July 23 to July 31, 2005, dust event (i), of (a) OMI tropospheric NO₂ concentrations and (b) MODIS-based PM_{2.5} to PM₁₀ ratio values. On the bottom: averaged values for October 4 to October 10, 2007 period, dust event (ii), of (c) OMI tropospheric NO₂ map and (d) MODIS-based PM_{2.5} to PM₁₀ ratio map.

Similarly, OMI-derived AOP presented in Fig. 7 is computed as averaged over the October 4 to October 10, 2007 time window. In particular, the SSA values averaged over Lake Garda result equal to 0.92 ± 0.02 , 0.93 ± 0.02 , and 0.96 ± 0.02 at 354, 388, and 500 nm, respectively [cf. Fig. 6(b), square yellow symbol], also providing a spectral increasing behavior for pixels where desert dust aerosol occurs. Corresponding EAOD_n data are shown in the left plot of the same figure [Fig. 6(a)]; these values are compatible with a presence of local aerosol (rural and/or urban) mixed with a dust-transported component.

Similarly to what has been performed in the analysis of the first dust event, Fig. 8 presents satellite-based information used to provide an exhaustive picture of the second event. Thus, NO₂ tropospheric column amount is shown [Fig. 8(c)] as derived from OMI observations over the October 4 to October 10, 2007 period. The geographical distribution of the NO₂ columnar content reflects the well-known very high pollution due to anthropogenic activities of the area.

Together with this, the MODIS-based PM_{2.5}/PM₁₀ ratio [Fig. 8(d)] in the Lake Garda area provided values around 0.6 and 0.7, highlighting the presence of coarse particles and thus confirming the presence of dust transported in this area.

The two no-dust case studies, events (iii) and (iv) (cf. Table 1), are presented with the aim of providing a benchmark for the comparison with the dust events singled out, confirming the capability of the satellite data to reliably detect and monitor the desert dust transport from the sources toward the domain of interest.

The OMI AOP spectral data relative to Lake Garda area and related to the first half of June 2005—event (iii)—are represented in Fig. 6. In this case, the retrieved EAOD_n and SSA data (square gray symbols) have been averaged between June 1 and June 14, 2005 for the three OMI aerosol wavelengths. The spectral trend displayed by these parameters is characteristic of a

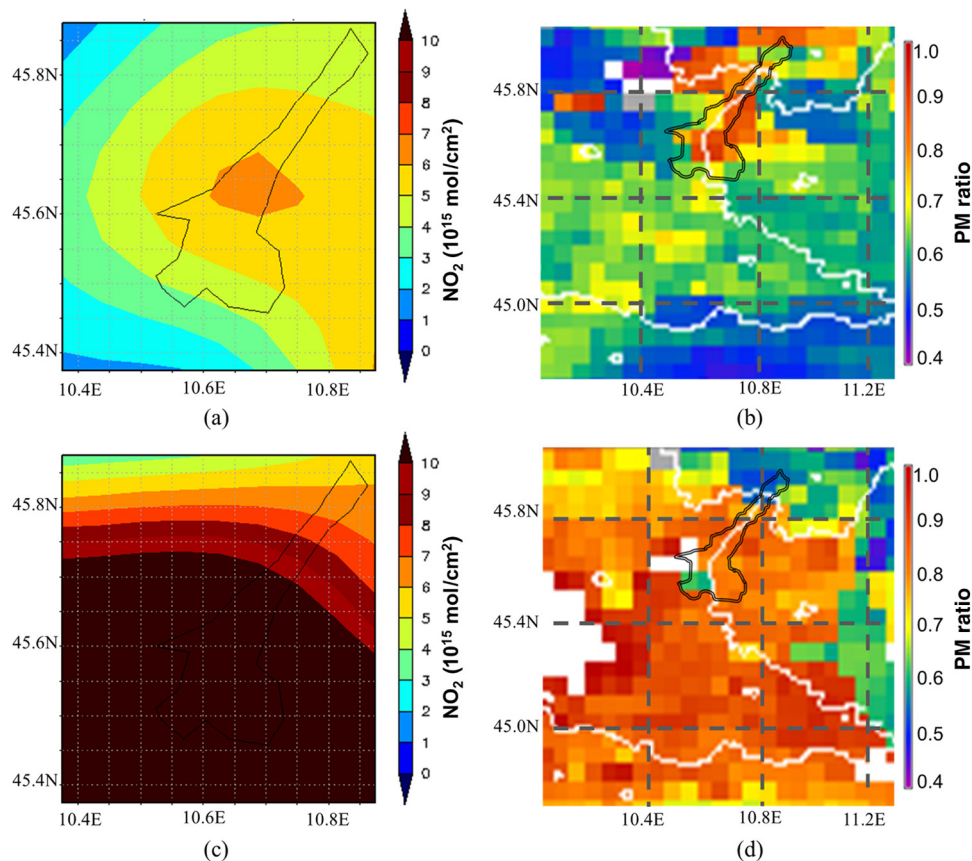


Fig. 9 On the top: average of period June 1 to June 15, 2005, no-dust event (iii), of (a) OMI tropospheric NO₂ concentrations and (b) MODIS-based PM_{2.5} to PM₁₀ ratio values. On the bottom: averaged values for October 30 to November 30, 2006 period, no-dust event (iv), of (c) OMI tropospheric NO₂ map and (d) MODIS-based PM_{2.5} to PM₁₀ ratio map.

Table 2 Atmospheric parameters derived from OMI/Aura (EAOD_n^{*}, SSA, NO₂, UV aerosol index) and MODIS/Terra and Aqua (PM_{2.5} to PM₁₀ ratio) averaged over time period of the events indicated in Table 1 for the Lake Garda area. The uncertainties represent the spatial variability of averaged value (i.e., the RMS linked to the mean spatial value). The percentage of mixing ratio by number [C_m , see Eq. (2)] of the modeling aerosol type is also reported for each Lake Garda event.

| Satellite-retrieved parameter | Lake Garda—event | | | |
|--|-------------------------|-------------|---------------|---------------|
| | (i) | (ii) | (iii) | (iv) |
| EAOD _n [*] [] at 354 nm | 1.26 ± 0.10 | 1.29 ± 0.10 | 1.57 ± 0.05 | 1.65 ± 0.02 |
| EAOD _n [*] [] at 354 nm | 1.14 ± 0.10 | 1.18 ± 0.10 | 1.32 ± 0.05 | 1.37 ± 0.02 |
| SSA [*] [] at 354 nm | 0.96 ± 0.01 | 0.92 ± 0.01 | 0.965 ± 0.005 | 0.970 ± 0.005 |
| SSA [*] [] at 388 nm | 0.97 ± 0.01 | 0.93 ± 0.01 | 0.960 ± 0.005 | 0.965 ± 0.005 |
| SSA [*] [] at 500 nm | 0.99 ± 0.01 | 0.96 ± 0.01 | 0.955 ± 0.005 | 0.955 ± 0.005 |
| PM _{2.5} /PM ₁₀ [] | 0.61 ± 0.05 | 0.64 ± 0.08 | 0.83 ± 0.03 | 0.88 ± 0.02 |
| NO ₂ [10 ⁵ mol/cm ²] | 2.6 ± 0.5 | 8.8 ± 1.0 | 5.5 ± 1.5 | 10.1 ± 1.0 |
| UV aerosol index [] | 0.8 ± 0.2 | 1.0 ± 0.4 | 0.8 ± 0.1 | 0.9 ± 0.1 |
| Aerosol type | Mixing ratio, C_m (%) | | | |
| Desert dust CV | 89 | 75 | 1 | 2 |
| Rural/clean continental | 8 | 19 | 61 | 28 |
| Urban | 1 | 5 | 35 | 67 |
| Other | 2 | 1 | 3 | 3 |

scattering aerosol polydispersion suspended in the troposphere with the spectral extinction (EAOD_n) decreasing and typical of a mixing of urban (black circle) and mainly rural (green circle) aerosols components. In particular, the OMI-retrieved SSA values averaged over Lake Garda slightly decrease with wavelength with a value of 0.96 ± 0.02 at the three OMI wavelengths and are very similar to SSA values of rural/clean continental aerosol model (values around 0.95 of green circles).

NO₂ vertical column [Fig. 9(a)] and MODIS-based PM_{2.5}/PM₁₀ ratio [Fig. 9(b)] highlight the presence of anthropogenic pollution in the analyzed period of June 2005. In particular, the PM_{2.5}/PM₁₀ ratio over Lake Garda was around 0.7 and 0.8.

The month of November 2006—event (iv)—has been selected as the second case study for a no-dust analysis. Similarly to the previous cases, Fig. 6 shows the averaged values of the period for Lake Garda of EAOD_n and SSA (black squares) at the three wavelengths. Again, the spectral trend of the parameters confirms the presence in the troposphere of a basically scattering aerosol polydispersion, with a significant presence of urban aerosol type (black circles in Fig. 6). The satellite-based observations in Fig. 9(c) show values of NO₂ columnar amount, highlighting a significant contribution of human activities to air pollution. The MODIS-based PM_{2.5}/PM₁₀ ratio [Fig. 9(d)] provides averaged values higher than 0.8, indicating a main presence of fine particles in the study area.

The first five rows in Table 2 report the values of the whole data set of the atmospheric parameters retrieved from OMI/Aura and MODIS/Terra and Aqua satellite sensors. All data are averaged over each temporal period indicated in Table 1 and over the Lake Garda area.

A further indicator of the desert dust aerosol suspended in the troposphere is represented by the aerosol index (AI) that can be considered as an indicator of the presence of UV-absorbing aerosols (as mineral dust and smoke) against nonabsorbing aerosols (as sea salt and sulfate particles).⁴⁷ Accordingly, last four rows in Table 2 show OMI AI for each Lake Garda—event.

Typically, AI values greater than 2 highlight the presence of dust. In the analyzed Lake Garda—events, the retrieved OMI AI values are between 0.8 and 1.0 also in the case of dust events (i) and (ii). Actually, these values prevent this parameter to be considered as a clear indicator of the presence of absorbing aerosol, like desert dust. This is mainly due to the small values of columnar EAOD. In fact, in the event (i) inside the dust plume over the Mediterranean area (Fig. 5) AI assumed values greater than 3.

To also provide a quantitative estimate of composition of the satellite-retrieved aerosol for each event, it can be assumed that the retrieved aerosol is composed by an external mixing³⁰ of the different aerosol types, desert CV, rural/clean continental, and urban, retrievable over the Lake Garda area. Thus, if C_m [0 ÷ 1] is the mixing ratio by the number of the m 'th aerosol type and $EAOD_n^*$ are the OMI-retrieved values reported in Table 2, the following system of equations can be considered

$$EAOD_n^*(\lambda) = \sum_{m=1,2,3} C_m EAOD_{n,m}(\lambda) \quad \text{for } \lambda = 354, 388, 500 \text{ nm}, \quad (2)$$

where $m = 1, 2, 3$ for desert CV, rural/clean continental, and urban, respectively; $EAOD_{n,m}$ represents the spectral values of the EAOD of the four aerosol models reported in Fig. 6. From Eq. (1), $EAOD_n^* = EAOD_n = 1$ at $\lambda = 500$ nm.

For each Lake Garda—event, the values of C_m are also reported in Table 2 as solutions of Eq. (2). In the last row of the table, the percentage of the $EAOD_n^*$ is also reported as residual of the right member of Eq. (2). This is the fraction of EAOD retrieved, which cannot be fully explained by Eq. (2). Solution of this equation confirms a significant presence of desert dust for the events (i) and (ii) (89% and 75%, respectively), while for the events (iii) and (iv), the most important components are the rural/clean continental (61%) and urban (67%), respectively.

3.2 Effect of Dust Deposition on Phytoplankton

The evaluation of the effect of dust deposition on phytoplankton abundance was carried out through the analysis of the two dust events (i) and (ii) and on the second no-dust event (iv). However, to evaluate the actual effect of dust deposition on phytoplankton growth, the following factors, which also might influence the phytoplankton dynamics, were considered. (i) The precipitation that influences wet deposition and allows the run-off of all material deposited along the coast was analyzed; while wet depositions have an impact on phytoplankton abundance,⁴⁸ dry depositions seem to have no significant effects. For this purpose, daily cumulate data from meteorological stations in Ponti S.M. and Bormio were evaluated. Other factors taken into account were (ii) the lake surface temperature (LST) and (iii) the occurrence of complete vertical circulation of the lake. In particular, these last two are both important factors affecting phytoplankton, whose primary production is correlated to LST trend^{49,50} and nutrients (as mainly nitrogen, phosphorus, and trace metals) whose availability in oligomictic deep lakes, such as Lake Garda, significantly increase in case of complete water column vertical circulation.

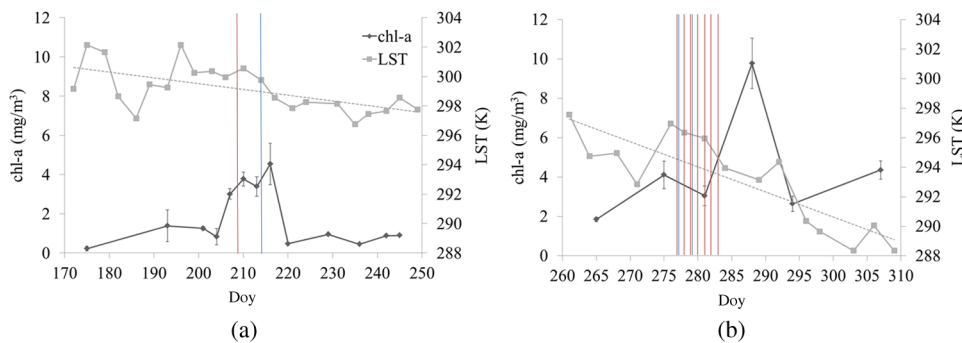


Fig. 10 Time pattern of chl-a concentration and LST on the time window of dust events in (a) July 2005 and (b) October 2007. Orange lines show dust event dates, blue lines show precipitation occurrences.

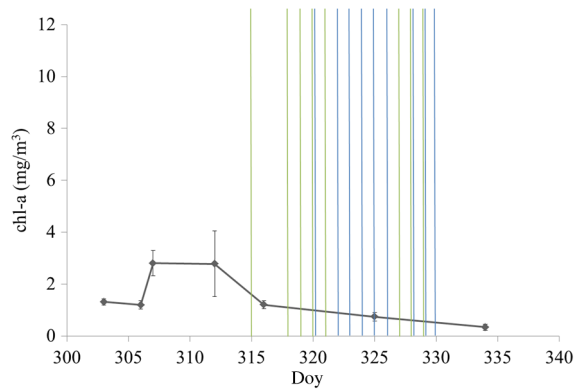


Fig. 11 Chl-a concentration trend in the time window of the no-dust event in November 2006 [event (iv) of Table 1]. Green lines show PM_{10} peaks dates, blue lines show precipitation occurrences.

The temporal dynamic of MERIS-derived chl-a concentrations in the pelagic area during the two dust transport events is shown in Figs. 10(a) and 10(b), together with dust event date and precipitation occurrence. In October 2007, precipitation was recorded simultaneously to dust transport event, allowing wet deposition, while in July 2005 no precipitation occurred until

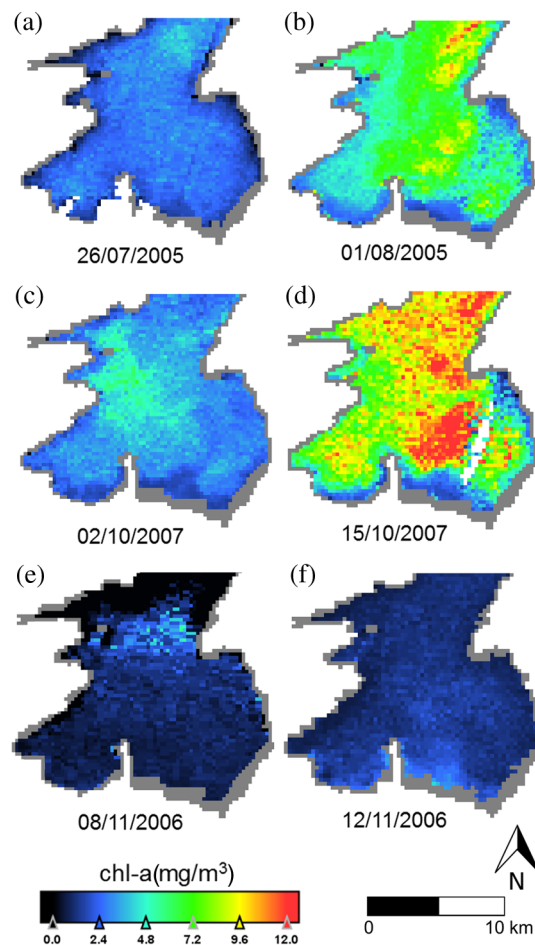


Fig. 12 Chl-a concentration before and after (a and b) dust event (i), (c and d) dust event (ii) and (e and f) no-dust event (iv). Gray indicates the pixels whose spectrum is too contaminated by the bottom signal and the adjacency effect; white indicates the invalid pixels (due to cloudiness or radiometric noise). The narrower northern part of the lake was omitted from the analysis due to the noise caused by the mountain adjacency.

6 days after dry deposition, this last confirmed by *in situ* PM measurements. However, an increase in chl-a concentration and thus in phytoplankton abundance (in the time window of desert dust transport) is clearly recognizable after both events; the percentage ratio increase in the chl-a concentration estimated between the last image before the event and the maximum value recorded after the deposition have been about 34% and 173%, respectively, for the case relative to July 28, 2005 and October 2007. The time-trend of LST showed that no increase occurred in the evaluated time window but, indeed, a negative trend was recorded as shown in Figs. 10(a) and 10(b), suggesting the independence of chl-a concentration increase from LST in both cases.

The time-lag between dust wet deposition (or the first precipitation recorded after the dust dry deposition) and the date of the first increment and maximum concentration was calculated. In this case, the first increase and maximum coincide; it turned out to range between 2 days in July 2005 and 8 days in October 2007. These results are in line with those in literature: e.g., from Ref. 48, in dust experiment in a low nutrient, low chlorophyll ecosystem (DUNE) project chl-a concentration increase could be recorded within 24 h and lasted several days (at least a week).

A similar analysis was performed for the no-dust event on November 2006, for which MERIS images were available. In this case, as it can be seen in Fig. 11, no appreciable increase of chl-a has been recorded from MERIS.

Figure 12 shows chl-a maps retrieved for the last image before and the first image after each event considered; an increase of chl-a concentration affects the whole lake area shown in the map after dust events, excepted for the coastal area. In fact, the effect is greater moving away from the coastal area where other factors lead chl-a concentration dynamics (e.g., anthropogenic discharges).

4 Conclusions

The findings of the analysis presented in this work outlined the capabilities of multispectral and multisensors satellite data to trace desert aerosol transport and the response of phytoplankton in lake waters.

A selection of case studies relative to desert dust transport episodes in Northern Italy have been singled out based on ground observations gathered from PM and AERONET stations. For these events, remote sensing images from SEVIRI/MSG, OMI/Aura, MODIS/Terra-Aqua, and MERIS/Envisat have been analyzed for (i) evaluating the dust transport and deposition in the Lake Garda area and (ii) investigating the effects of the dust presence on phytoplankton growth in lake waters.

Together with these cases, periods characterized by absence of desert dust aerosol, the no-dust events, have also been selected. This has been done to allow the comparison with the dust events and thus confirming the capability of the satellite remote sensing techniques to reliably detect this transport from the sources toward the domain of interest. In particular, SEVIRI/MSG observations have been used to detect and follow desert dust transport pattern on the basis of the differences between BTs recorded in various IR channels. Measurements in the UV–VIS range by OMI/AURA have been analyzed providing estimates of spectral AOP and the percentage of the external mixing ratio by the number of desert dust, rural, and urban aerosol types. Together with these satellite data, MODIS-derived $PM_{2.5}/PM_{10}$ ratio maps for the four analyzed periods confirmed the presence (or the absence) of desert dust in the study area, being the ratio values on average less (or greater) than about 0.65.

After that, MERIS-derived chl-a concentrations have been used for the evaluation of the dust deposition effect in Lake Garda. Results showed an increase of chl-a concentration with values varying from 30% to 170% and a time-lag of about 2 to 8 days. In occasion of the no-dust event, the same techniques turned out to provide no significant variation in chl-a concentration. Furthermore, a large set of *in situ* data was employed to fully describe and confirm the effect of dust deposition. Thus, the wide set of observations (both satellite and ground-based) and related techniques emphasizes how synergic approach might provide an exhaustive picture of the desert dust case studies and of the subsequent effect on lake water quality.

Nonetheless, new case studies have to be considered to further investigate the response of the lake to dust deposition events. In this view, it will be possible to go ahead by exploiting the

Table 3 Summary of the main features of forthcoming satellite sensors in the frame of the next EO missions. The spatial resolution is relative to subsatellite point.

| Satellite sensor platform | Purpose | Features | Spatial resolution | Expected lifetime | Satellite sensor heritage |
|---------------------------|--|--|------------------------|-------------------|---------------------------|
| VIIRS SUOMI-NPP | Atmosphere, land ocean | 22 channels VIR/IR radiometer | 275 to 750 m | Up to 2029 | MODIS |
| TROPOMI Sentinel 5P | Atmospheric chemistry: trace gases and aerosol | UV–VIS grating spectrometer | 7 km | 2014 to 2021 | OMI |
| MSI Sentinel-2A, 2B | Ocean color, vegetation, aerosol, clouds | 13 channels from VNIR to SWIR | 10 to 60 m | 2015 to 2023 | LANDSAT |
| OLCI Sentinel-3A, 3B | Ocean color, vegetation, aerosol | 21 channels VIS/NIR optical imager | 300 m | 2015 to 2023 | MERIS |
| SLSTR Sentinel-3A, 3B | Aerosol, clouds, surface | 9 channels VIR/IR optical imager | 0.5 km 1 km for TIR | 2015 to 2023 | AATSR |
| UVN Sentinel-4 | Atmospheric chemistry: trace gases and aerosol | UV–VIS NIR grating spectrometer | <8 km | 2021 to 2029 | OMI |
| FCI MTG-I | Clouds, water vapor features | 16 channels VIS, NIR, SWIR, MWIR, and TIR | 0.5 to 2 km | 2019 to 2027 | SEVIRI |

observations that will be carried out from the forthcoming satellite sensors, within the frame of next EO missions. Actually these planned activities will allow providing a significant improved monitoring of the atmospheric constituents, mainly in terms of increasing frequency of observations, giving the unique possibility for an operational monitoring of the environmental impacts of desert dust transport and deposition. In addition, they will allow improvement of studies on effects on phytoplankton abundance providing new suitable sensors for water quality retrieval such as OLCI, which will be the successor of MERIS (presenting similar spatial and spectral resolutions) and will be onboard the two Sentinel 3A and 3B.

Table 3 presents a list of recently launched and forthcoming satellite sensors that could be used in this kind of study. Each sensor is presented with the corresponding hosting platform, and then the main purpose for which the sensor is intended follows together with its characteristic of spectral and spatial resolution and its expected lifetime. Finally, in the table, the heritage, on the basis of which the new sensor is built, is given in order to highlight the continuity with the old or in operation sensors used in this work.

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